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IV-2 ONE-DIMENSIONAL FLOW

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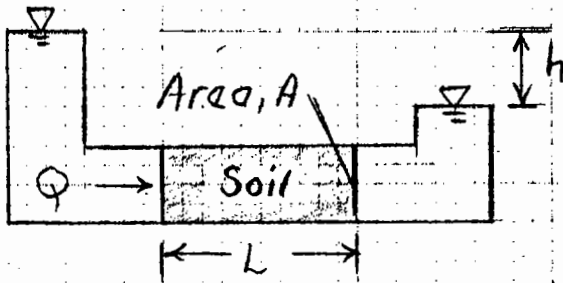
13 787
 500 SHEETS MILLER 5 SQUARE
 600 SHEETS MILLER 5 SQUARE
 100 SHEETS EYM-EASE 5 SQUARE
 42-382 200 SHEETS EYM-EASE 5 SQUARE
 42-389 200 SHEETS EYM-EASE 5 SQUARE
 42-392 200 RECYCLED WHITE 5 SQUARE
 42-395 200 RECYCLED WHITE 5 SQUARE
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Part II - 2 ONE DIMENSIONAL FLOW (Chap. 17)

1. FUNDAMENTALS OF FLOW

1.1 Darcy's Law (Laminar flow)



Flow Q (L^3/t)
 $= k \frac{h}{L} A = \underline{\underline{kiA = Q}}$

where

k = coef. of permeability (L/D)^{*}

h = change in total head (loss thru soil) (L)

i = hydraulic gradient = $\frac{h}{L}$ (dimensionless)

* Now called hydraulic conductivity

Note: Turbulent flow if high i + coarse grained

1.2 Heads (= energy / "unit weight")

(1) general

Total head $h = h_{\text{physical-chemical}} + h_{\text{velocity}} + h_{\text{elevation}} + h_{\text{pressure}}$

Causes Flow Can neglect Potential Energy

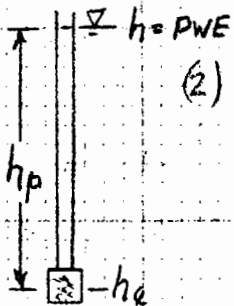
force/unit volume $\frac{1}{2} \frac{v^2}{g}$ } Kinetic Energy

Except 1st for contaminant transport

(2) Used for flow thru soils

$h = h_e$ (elevation head wrt some arbitrary datum)
 $+ h_p$ (pressure head = u/γ_w = height of H_2O in piezometer)

Piezometric water elevation (PWE) = TOTAL HEAD = elevation of top of water in piezometer



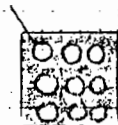
1.3 Flow Velocities

(1) Approach velocity (v)

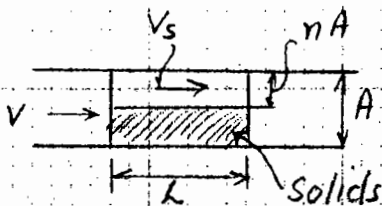
Area A
 $Q \rightarrow \text{Soil} \rightarrow v = \frac{Q}{A} = \underline{\underline{ki = v}}$

$n = \frac{V_v}{V} = \frac{A_v}{A}$

$\Sigma V = \text{Area Voids} = n \text{ Area}$



(2) Seepage velocity (v_s)



If treat soil as straight capillary tubes
 Volume Voids = n Volume soil

$Q = VA = v_s A_v = v_s nA$ ($\frac{L \times A_v}{L \times A} = \frac{V_v}{V} = n$)

$\therefore v_s = \frac{v}{n}$ and $t = \frac{L}{v_s} = \frac{Ln}{v}$

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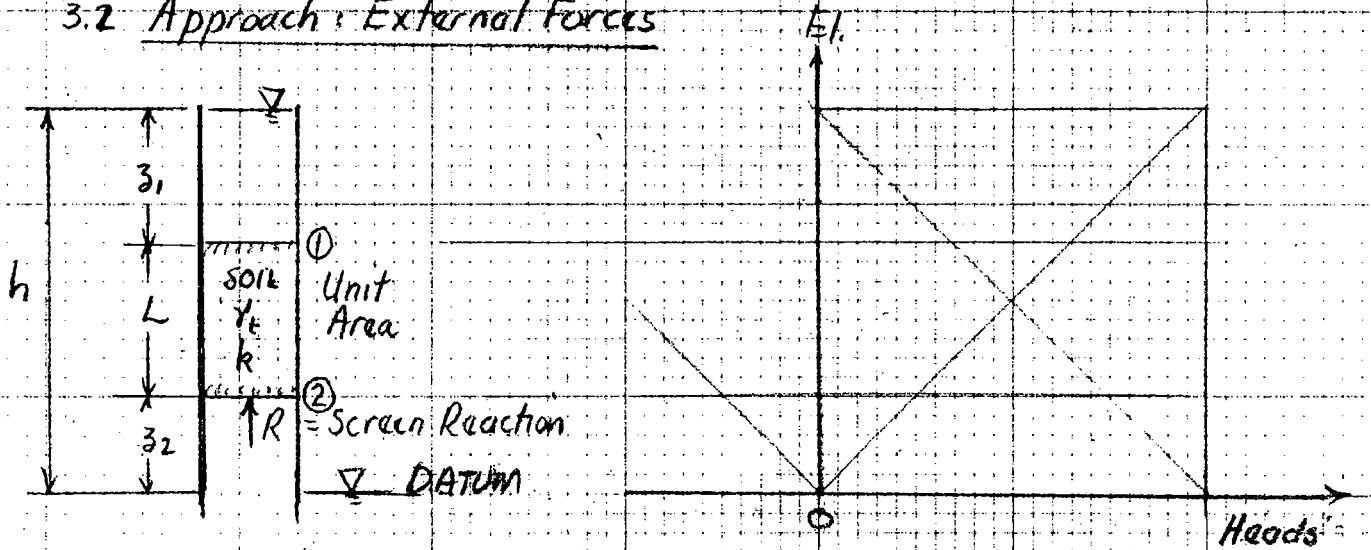
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3. COMPUTATION OF EFFECTIVE STRESSES WITH SEEPAGE

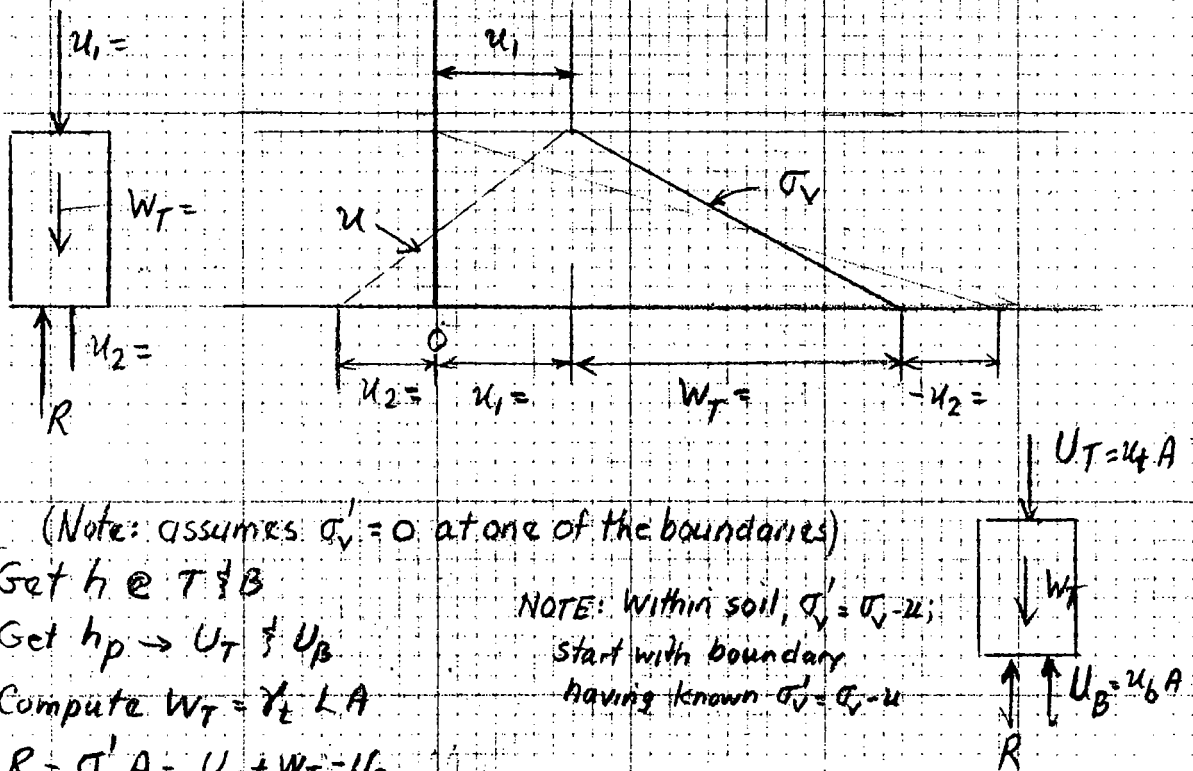
3.1 Two Approaches (For both, always get $h = h_e + h_p$ for "system")

- "External Forces" = Total Weights + Boundary (Water) Forces
- "Internal Forces" = Buoyant Weights + Seepage Forces

3.2 Approach: External Forces



Isolate Soil (Unit area)



Steps: (Note: assumes $\sigma'_v = 0$ at one of the boundaries)

- (1) Get h_e @ T & B
- (2) Get $h_p \rightarrow U_T \& U_B$
- (3) Compute $W_T = \gamma_s L A$
- (4) $R = \sigma'_v A = U_T + W_T - U_B$

NOTE: Within soil, $\sigma'_v = \sigma_v - u$; start with boundary having known $\sigma'_v = \sigma_v - u$

NOTE: R can act \uparrow or \downarrow (also can have R forces acting on both boundaries)

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3.3 Approach: Internal Forces

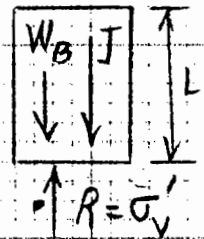
- Reaction = force due to buoyant weight of soil ($W_B = \text{magnitude w/o flow} = \gamma_b L$) + drag forces due to seepage ($J = jL$)

- Define $j = \frac{\text{seepage force}}{\text{unit volume}} = i \gamma_w = \frac{h}{L} \gamma_w$ which acts in direction of flow

- Comparison of 2 approaches

$R =$

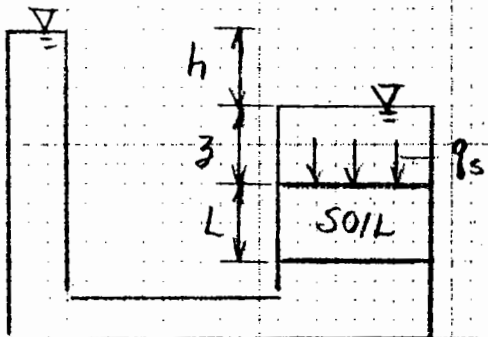
Unit Area



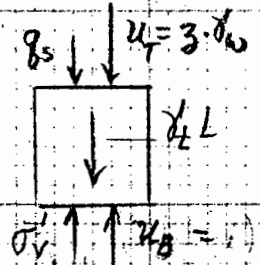
$$R = W_B + J = \gamma_b L + \frac{h}{L} \gamma_w L \rightarrow h \gamma_w$$

NOTE: If $\gamma_f \neq \gamma_w$, then $\gamma_b = \gamma_s - \gamma_f$ and $j = i \gamma_f$

3.4 Example With Upward Flow



Unit Area



(1) External Forces

$$R = \sigma'_v = q_s + \gamma_s L + jL - (h + z + L) \gamma_w = q_s + L(\gamma_s - \gamma_w) - h \gamma_w$$

(2) Internal Forces: $\sigma'_v = q_s + \gamma_b L - J \left\{ = jL = i \gamma_w L = \frac{h}{L} \gamma_w L = h \gamma_w \right\}$

(3) For $R = \sigma'_v = 0$ at bottom

$$h_{cr} = \frac{q_s + \gamma_b L}{\gamma_w}$$

(4) For $q_s = 0$ ($\therefore \sigma'_v = 0$ throughout)

$$h_{cr} = \frac{\gamma_b L}{\gamma_w} \text{ or critical gradient } i_{cr} = \frac{\gamma_b}{\gamma_w}$$

3.5 "Quick" Condition (Quick sand when $\sigma'_v = 0$ cohesionless soil)

- Acts as heavy fluid since strength $\rightarrow 0$ when $\sigma'_v \rightarrow 0$
- $i_{cr} = 1.04 \pm 0.08$ for sandy materials, and independent of soil thickness ($\gamma_b = 65 \pm 5 \text{ pcf}$)
- Fine sands & cohesionless silts most susceptible as require less flow
- Reality vs movies